

## Timing of thermal episodes

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The two previous chapters have demonstrated a consistent pattern of thermal maturity and a distribution of bitumen occurrences suggesting that significant amounts of hydrocarbons were once generated and expelled with possible later migration. The timing of this migration has important economic aspects, especially in relation to the formation of traps and the degradation of accumulated hydrocarbons.

Precise dating of hydrocarbon generation, migration and accumulation is difficult unless these processes are clearly related to specific geological events such as: (1) formation of diagenetic minerals, which may be dated isotopically (e.g. smectite, illite, glauconite); or (2) magmatic events which may be dated isotopically (e.g. granite or dolerite intrusion). Indirect evidence, often producing relatively narrow time constraints, may be obtained from: (1) the subsidence history deduced from stratigraphic and sedimentological data; (2) drastic changes in thermal maturity or maturity patterns across unconformities; or (3) the thermal history (especially uplift) deduced from fission track age determinations and fission track length distributions.

### Franklinian subsidence and Ellesmerian orogeny

Stratigraphic and sedimentological knowledge of the Franklinian basin in North Greenland, recently described in detail by Higgins *et al.* (in press), provides a good background for the consideration of the Early Palaeozoic subsidence history and its thermal maturity effect on the source rocks.

Shelf carbonates deposited along the southern margin of the Franklinian basin seem particularly suitable for calculation of the subsidence history. Preliminary determinations by M. S nderholm (GGU, unpublished) suggest slow steady subsidence with a slightly decreasing rate of 15–20 mm/1000 y from Early Cambrian to Early Silurian time. Estimated rates for the Late Silurian indicate much higher values around 125–150 mm/1000 y.

The sediments in the northern part of the region were laid down in a deep-water trough and indicate considerably higher subsidence rates throughout the Early Palaeozoic (especially in the Early Cambrian), apparently also with increasing values in the Silurian.

Little is known of possible continued subsidence as Devonian sediments are known only in the Canadian part of the Franklinian basin. However, it is clear that the Cambrian source rocks in the northernmost areas (north of the Navarana Fjord Escarpment), and probably also some of the Silurian shales, were already buried so deeply by Late Silurian or Early Devonian time that hydrocarbon generation took place. In the central and southern parts of the region it is not known whether the burial history can account for such an early generation of hydrocarbons.

The Late Palaeozoic Ellesmerian orogeny brought sedimentation in the Franklinian basin to a close (Trettin & Balkwill, 1979). In North Greenland it gave rise to the North Greenland fold belt, which has an E–W trend across most of the northern part of the region and roughly coincides with the trend of the deep-water trough. The rocks affected by deformation are mainly turbiditic sandstones and shales of Cambrian to Early Silurian age, and exhibit an increasing intensity of deformation and metamorphism towards the north (see fig. 60; Dawes, 1976; Higgins *et al.*, 1982, 1985; Soper & Higgins, 1987).

The youngest known sediments affected by the Ellesmerian orogeny in North Greenland are of latest Silurian, possibly earliest Devonian age (Peel, 1982; Higgins *et al.*, in press), whereas the oldest post-Ellesmerian sediments which unconformably overlie folded rocks are of Late Carboniferous age (H kansson & Stemmerik, 1984; Stemmerik & H kansson, in press). Rb–Sr isotopic dating of whole rocks and clay mineral fractions suggests a Late Devonian to Early Carboniferous age for the Ellesmerian metamorphism (Springer, 1981; N. Springer & J. D. Friderichsen, personal communication, 1988).

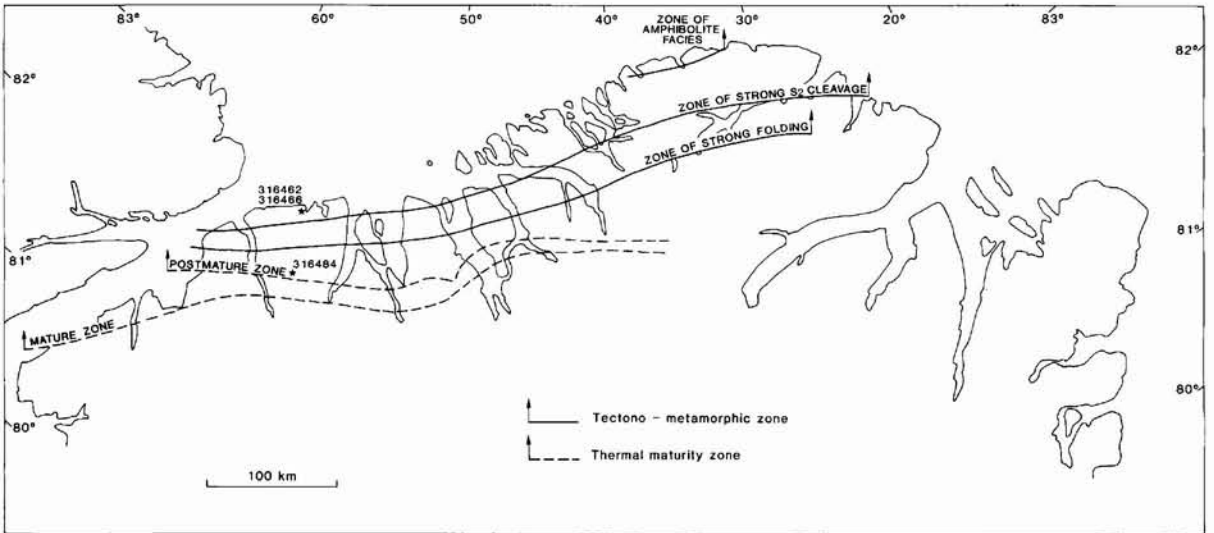


Fig. 60. Map of Ellesmerian tectono-metamorphic zones. Modified after Dawes (1976) and Higgins *et al.* (1982, 1985). The thermal maturity zonation is modified from Chapter 6. The locations of samples studied by the fission track method are indicated by stars.

The striking parallelism of the Ellesmerian tectono-metamorphic zones and the thermal maturity pattern (fig. 60) suggests that the strong thermal effect in the Early Palaeozoic source rocks is largely Ellesmerian. The thermal maturity increases markedly from thermally immature/early mature source rocks in the south to postmature source rocks over south–north distances of only 10–30 km as the fold belt is approached. A marked contrast in thermal maturity is observed in eastern Peary Land between thermally postmature or low metamorphic Silurian source rocks and overlying thermally immature or early mature Upper Palaeozoic and Mesozoic rocks (fig. 48; Christiansen *et al.*, in press). This implies at least two periods of burial or thermal events, one pre- or syn-Ellesmerian and one post-Ellesmerian, separated by a period of uplift.

### Late Palaeozoic and Mesozoic sedimentation

Late Palaeozoic and Mesozoic sediments are restricted in North Greenland to relatively small areas in the northern and eastern parts of Peary Land (fig. 61; Dawes, 1976; Håkansson & Stemmerik, 1984). They rest unconformably on older rocks, or occur as fault-bounded inliers. In western North Greenland Late Palaeozoic and Mesozoic sediments have not been preserved, and it is not known where they were deposited. However, rocks of this age are widely preserved in northern Ellesmere Island (Trettin & Balkwill, 1979).

The absence of post-Ellesmerian sediments in western North Greenland may indicate that the area was subjected to uplift during Late Palaeozoic and Mesozoic time, while deposition in the Peary Land region may have been confined to fault-controlled basins.

### Cretaceous–Tertiary magmatic and tectonic events

Cretaceous to Tertiary magmatic and tectonic activity is widespread in the eastern and northern part of North Greenland (fig. 61) (Soper *et al.*, 1982; Higgins *et al.*, 1985; Friderichsen & Bengaard, 1985). It is referred to the Eurekan orogeny, which affected the Canadian Arctic Islands as well as North Greenland (Trettin & Balkwill, 1972; Trettin *et al.*, 1979).

Dolerite dykes were intruded in dense N–S trending swarms along the north coast of Greenland, in north Peary Land and Nansen Land. Further south dykes are less frequent and have E–W to SE–NW trends (fig. 61). Geochemically the dykes are of alkaline type ('within-plate') (Soper *et al.*, 1982). The dykes cross-cut early Cretaceous sediments and are post-dated by the Kap Washington Group peralkaline volcanics and interbedded sediments which have a latest Cretaceous to earliest Tertiary age (Larsen *et al.*, 1978; Batten *et al.*, 1981; Batten, 1982; Larsen, 1982; Soper *et al.*, 1982). Friderichsen & Bengaard (1985) noted two dyke generations in Nansen Land, an early generation of densely distrib-

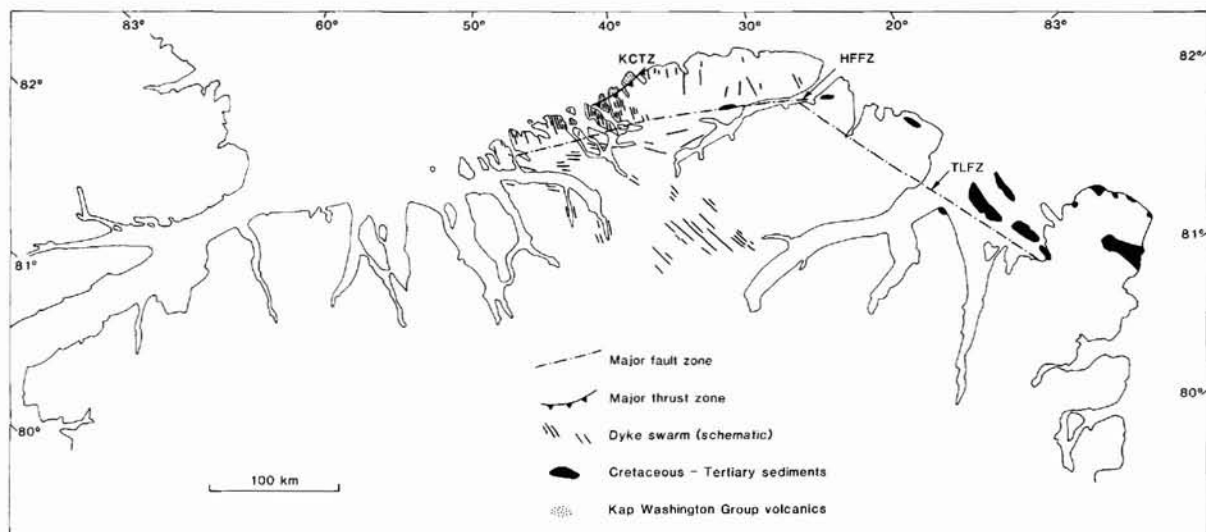


Fig. 61. Map of North Greenland showing Cretaceous-Tertiary magmatic and tectonic features. Modified after Soper *et al.* (1982). KCTZ: Kap Cannon thrust zone, HFFZ: Harder Fjord fault zone, TLFZ: Trolle Land fault zone.

uted N-S oriented dykes, and a later generation of scattered NW-SE to WNW-ESE oriented dykes.

Lower Palaeozoic metasediments of the North Greenland fold belt are thrust over the Kap Washington Group volcanics along the Kap Cannon thrust zone (fig. 61; Soper *et al.*, 1982). Isotopic ages suggest a Paleocene to Eocene age for the thrust event (Soper *et al.*, 1982).

Due to the limited occurrence of post-Ellesmerian sediments in North Greenland it is often difficult to distinguish Eurekan structures and events from those of Ellesmerian age. The most important Eurekan events are the northward thrusting on the Kap Cannon thrust zone and movements along the Harder Fjord and Trolle Land fault zones (fig. 61). In the Freuchen Land and inner J. P. Koch Fjord region, the scattered dolerite dykes are probably responsible for local anomalous thermal effect on the source rocks of the Henson Gletscher Formation (Chapter 6; Christiansen *et al.*, 1987).

### Fission track analysis

A total of 15 samples from two profiles through thermally immature to postmature areas were prepared for fission track analyses, but only three of these yielded apatite grains of the appropriate size fraction (see details of preparation and analytical technique in Chapter 3).

All three samples are from Nyeboe Land (fig. 60); samples 316462 and 316466 are from turbiditic sand-

stones in the upper part of the Early Silurian Merqujôq Formation, whereas sample 316484 is from a sandstone in the Late Silurian Wulff Land Formation. The two first samples come from a thermally postmature to low metamorphic area in the steeply inclined Nyeboe Land linear belt (Dawes, 1982) where a cleavage is developed in the interlayered mudstone beds. The third sample is from a thermally postmature area, apparently undeformed, situated 50 km to the south (fig. 60).

The two northernmost samples, 316462 and 316466, show similar apatite fission track ages (Table 5). The

Table 5. Fission track ages of apatite and zircon

Sample	Ma $\pm 1 \sigma$	No. grains	Ns	Ni	$\chi^2(P\%)$
<i>Apatite fission track ages</i>					
316462	206.79 $\pm$ 22.14	9	253	165	96.4
316466	207.18 $\pm$ 21.53	9	275	179	89.2
316484	224.45 $\pm$ 17.24	16	435	463	99.2
<i>Zircon fission track ages</i>					
316462*	457.73 $\pm$ 45.40	8	552	125	4.9
	411.11 $\pm$ 41.41	8	528	130	7.3
316484	347.74 $\pm$ 34.30	9	482	141	88.8

\*Two determinations; the very low probability of the chi-squared test ( $\chi^2(P\%)$ ), indicates that the interpretation of this age is difficult.

mean track lengths and track length distributions fall in the field of mixed and bimodal distributions of Gleadow *et al.* (1986) indicating a complex thermal history (Table 6; fig. 62). The bimodal character of 316462 and the broad based character of 316466 may be ascribed to a

Table 6. Confined track lengths of apatites

Sample	$t_{\text{mean}} \pm 1 \sigma$ ( $\mu\text{m}$ )	No.	Type	Quality
316462	$10.62 \pm 3.11$	35	bimodal	?good
316466	$9.01 \pm 3.19$	15	mixed	poor
316484	$10.99 \pm 2.53$	130	mixed	good

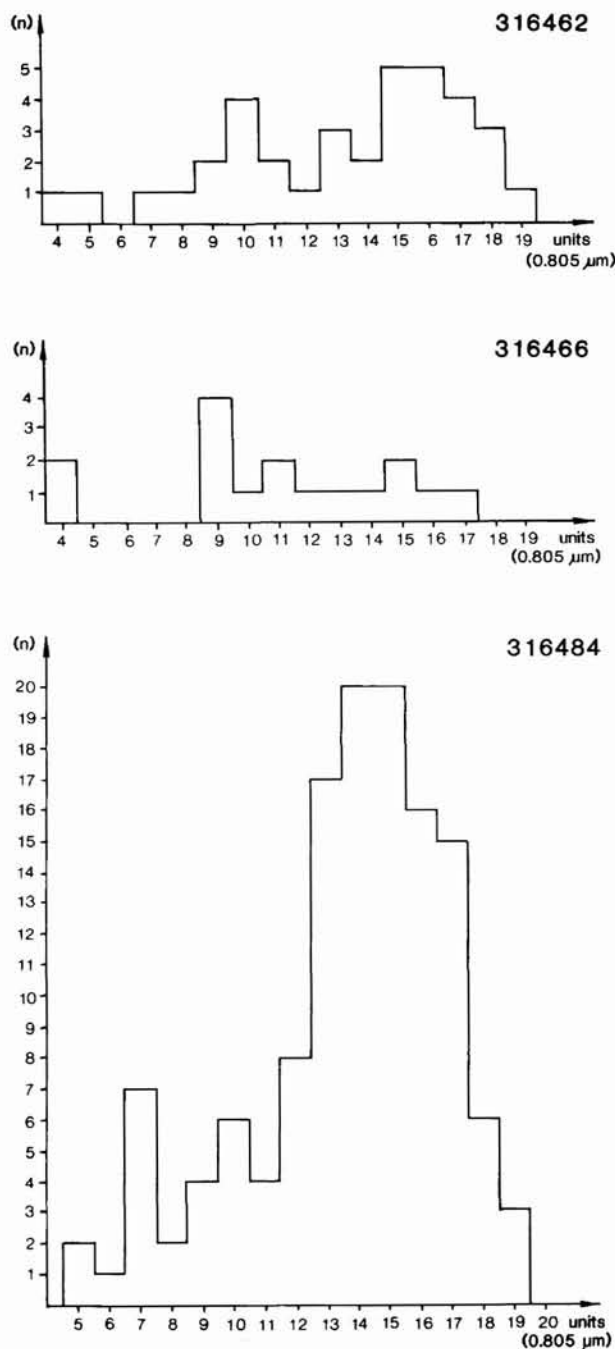


Fig. 62. Fission track length distributions. See details in Tables 5 and 6.

heating event in the 70–130°C range (Gleadow *et al.*, 1983) which only partly annealed earlier tracks. This event took place later than the age of 207 Ma. The southern sample, 316484, shows a slightly older age of 224 Ma and longer mean track length indicating a very limited late thermal overprint. This fission track age is expected to reflect the age of uplift through the 100°C isotherm.

Fission track zircon age determinations were performed on two of the samples (Table 5). The obtained fission track ages are significantly different from each other. The northern sample, 316462, was reanalysed, but in both cases gave a low probability of the chi-squared test. The individual zircon grain ages of the sample are close to, or older than, the time of deposition (c. 280–700 Ma) and may imply that pre-sedimentary fission tracks were only partly annealed in the temperature interval around  $240 \pm 50^\circ\text{C}$  (Hurford, 1986). The southern sample, 316484, is of much better analytical quality as expressed by the high probability of the chi-squared test, and the fission track age of 348 Ma obtained probably corresponds to the time of uplift through the 240°C isotherm.

A possible time-temperature history of the analysed samples is illustrated in fig. 63. All three sandstones were deposited in the Silurian and subjected to rapid subsidence shortly after deposition. During this subsidence and the following Ellesmerian orogeny the sandstones reached temperatures well above the apatite annealing range ( $100 \pm 30^\circ\text{C}$ ) and possibly also close to the zircon annealing temperature ( $240 \pm 50^\circ\text{C}$ ). Uplift through the 240°C isotherm took place shortly after the Ellesmerian orogeny, whereas uplift through the 100°C isotherm occurred in Late Triassic to Early Jurassic time. A possible late thermal overprint in the northern area was probably associated with Cretaceous to Tertiary magmatic activity.

Calculated cooling rates for sample 316484 give 8.9°C/Ma in the 240°C to 100°C interval and 4.5°C/Ma in the 100°C to 0°C interval. Using a geothermal gradient of 30°C/km, these values correspond to uplift rates of 36 m/Ma and 15 m/Ma, respectively. Such rates are comparable with post-Caledonian vertical movements in the

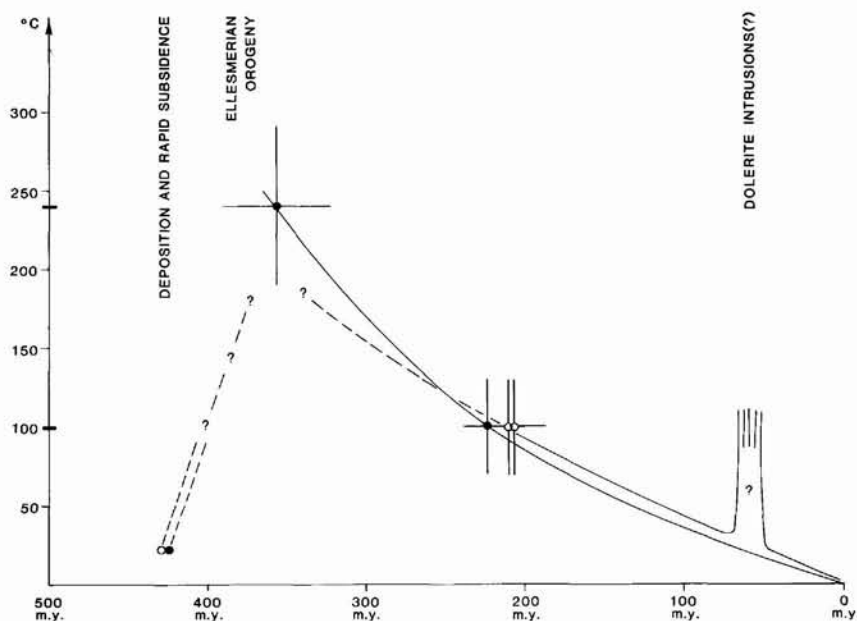


Fig. 63. Synoptic diagram showing the thermal tectonic history of the Silurian sediments in northern Nyeboe Land.

Scoresby Sund region of East Greenland (Hansen, 1985).

## Implications

The thermal history as outlined by geological considerations, the thermal maturity pattern, and the age determinations of thermal episodes and uplift, is consistent with a simple model of hydrocarbon generation and possible accumulation. Both the Middle to Late Cam-

brian and the Early Silurian source rocks were subjected to rapid subsidence in the Late Silurian and suffered a strong thermal alteration during this subsidence and the subsequent Ellesmerian deformation. Most of the potential hydrocarbons were probably already generated at this time. After the Ellesmerian orogeny, uplift and cooling prevailed throughout most of the Mesozoic, possibly with local heating in the northernmost part of the region in the Cretaceous and Tertiary in association with the intense dyke intrusion.



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## Plate 1. Kerogen

- A. Sample with a relatively low (0.21% TOC) kerogen content and a dominance of finely disseminated amorphous kerogen in a silica gel, Lafayette Bugt Formation, Nyeboe Land, GGU 316490-1, unsieved organic material.
- B. As A., Lafayette Bugt Formation, Nyeboe Land, GGU 316490-2, sieved organic material (on 10  $\mu\text{m}$  nylon mesh).
- C. Sample with a relatively moderate (1.15% TOC) kerogen content and small to moderate amounts of large amorphous kerogen particles, Thors Fjord Member, Nares Land, GGU 318007-18-1 unsieved organic material.

- D. As C., Thors Fjord Member, Nares Land, GGU 318007-18-2, sieved organic material (on 10  $\mu\text{m}$  nylon mesh).
- E. Sample with a relatively large (5.09% TOC) kerogen content and a dominance of large amorphous kerogen particles, Thors Fjord Member, Nares Land, GGU 318007-32-1, unsieved organic material.
- F. As E., Thors Fjord Member, Nares Land, GGU 318007-32-2, sieved organic material (on 10  $\mu\text{m}$  nylon mesh).

Scale bar: 20  $\mu\text{m}$ .

## Plate 2. Cambrian palynomorphs

- A. Acritarch-like folded alga. Middle Cambrian Sydpasset Formation, Freuchen Land, MGUH 19334 from GGU 315873-2; 139.5-13.9.
- B. Acritarch-like folded alga. Middle Cambrian Sydpasset Formation, Freuchen Land, MGUH 19335 from GGU 315873-2; 150.6-7.0.
- C. Two acritarch-like folded algae. Middle Cambrian Sydpasset Formation, Freuchen Land, MGUH 19336 (large light body), MGUH 19337 (dark small body), both from GGU 315873-2; 122.0-15.3.
- D. Acritarch-like folded alga. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19338 from GGU 324217-2; 139.2-14.9.
- E. Acritarch-like folded alga. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19339 from GGU 324300-2; 131.2-13.4.

- F. Acritarch-like folded alga. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19340 from GGU 324217-2; 141.9-11.4.
- G. Lump of algal or spore-like elements. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19341 from GGU 324300-2; 127.8-14.4.
- H. Diad-like lump of algal or spore-like elements. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19342 from GGU 314300-2; 138.5-17.8.
- I. Lump of alga or spore-like elements. Middle Cambrian, Ekspedition Bræ Formation, Freuchen Land, MGUH 19343 from GGU 324300-2; 157.7-14.7.

Scale bar: 20  $\mu\text{m}$ .

## Plate 3. Ordovician palynomorphs

- A. Acritarch. Upper Ordovician Troedsson Cliff Member, Washington Land, MGUH 19344 from GGU 316968-2; 145.3-17.8.
- B. Acritarch. Upper Ordovician - Lower Silurian Aleqatsiaq Fjord Formation, Washington Land, MGUH 19345 from GGU 316085-4; 124.1-21.3.
- C. Graptolite fragment, Upper Ordovician Troedsson Cliff Member, Washington Land, MGUH 19346 from GGU 316968-2; 128.1-2.9.
- D. Scolecodont, Upper Ordovician, Troedsson Cliff Member, Washington Land, MGUH 19347 from GGU 316968-2; 148.8-15.1.
- E. Alga. Upper Ordovician - Lower Silurian Aleqatsiaq Formation, Nyeboe Land, MGUH 19348 from GGU 316103-2; 135.0-15.0.
- F. Filamentous alga. Upper Ordovician - Lower Silurian Aleqatsiaq Formation, Washington Land, MGUH 19349 from GGU 316058-2; 135.1-4.2.

- G.-L. Spores with trilete rays. Upper Ordovician, Troedsson Cliff Member, Washington Land (Nøhr-Hansen & Koppellus, 1988).

G.-I. *Besselia nunaatica*, MGUH 17539 from GGU 316968-2; 125.5-8.3.

G. Distal view illustrating the minute ornamentation.

H. Equatorial view.

I. Internal proximal view.

J. *Besselia nunaatica*, two connected spores, internal proximal view, MGUH 17541 from GGU 316968-2; 155.1-11.9.

K.-L. *Besselia nunaatica*, MGUH 17542 from GGU 316968-2; 123.8-15.9.

K. Distal view illustrating the ornamentation.

L. Internal proximal view.

Scale bar: 20  $\mu\text{m}$ .

## Plate 4. Silurian palynomorphs

- A. Chitinozoan, *Angochitina* cf. *A. elongata*. Upper Silurian Wulff Land Formation, Wulff Land, MGUH 19350 from GGU 315950-3; 136.9-17.2.
- B. Chitinozoans, *Linochitina erratica*. Upper Silurian Wulff Land Formation, Wulff Land, MGUH 19351 from GGU 315950-2; 154.3-9.6.
- C. *Retiolites*, graptolite fragment. Upper Silurian, Wulff Land Formation, Wulff Land, MGUH 19352 from GGU 315950-3; 155.1-11.1.
- D. Graptolite fragment, Upper Silurian, Wulff Land Formation, Wulff Land, MGUH 19353 from GGU 315950-2; 127.3-5.7.
- E.-H. Trilete spore-like bodies, figs E and F with a degraded bitumen-like appearance.
- E. Lower Silurian Lafayette Bugt Formation, Washington Land, MGUH 19354 from GGU 211760-2; 143.3-17.2.
- F. Upper Silurian Wulff Land Formation, Wulff Land, MGUH 19355 from GGU 315950-3; 15950-3; 155.5-8.2.

G. Upper Silurian Nyeboe Land Formation, Nyeboe Land, MGUH 19356 from GGU 319234-2; 119.3-11.0.

H. Upper Silurian Nyeboe Land Formation, Wulff Land, MGUH 19357 from GGU 319210-3; 130.6-21.4.

I. Spherical folded algae, acritarchs? Lower Silurian Lafayette Bugt Formation, Hall Land, MGUH 19358 from GGU 324157-2; 144.2-8.5.

J. Tubular structure. Upper Silurian Nyeboe Land Formation, Wulff Land, MGUH 19359 from GGU 319210-3; 146.6-16.5.

K. Tubular structure. Upper Silurian Nyeboe Land Formation, Nyeboe Land, MGUH 19360 from GGU 319234-2; 138.8-8.0.

L. Rounded drop-shaped palynomorphs. Lower Silurian Lafayette Bugt Formation, Washington Land, MGUH 19361 from GGU 316061-2; 137.1-14.8.

Scale bar: 20  $\mu\text{m}$ .

Plate 1. Kerogen

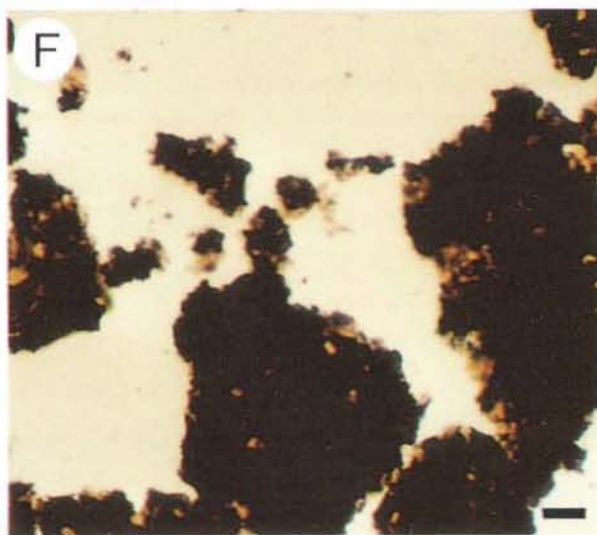
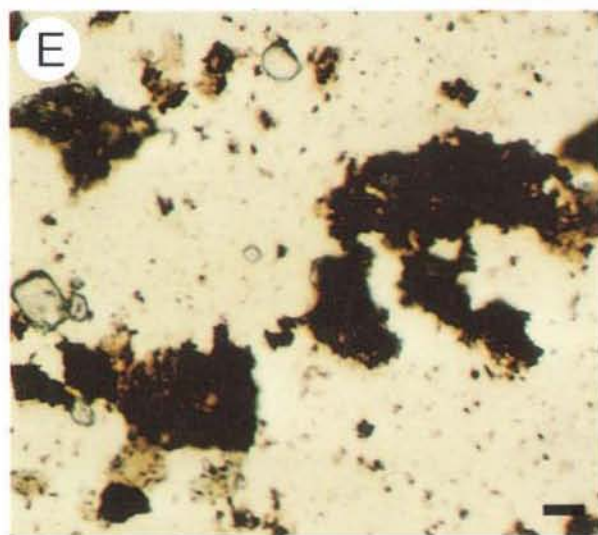
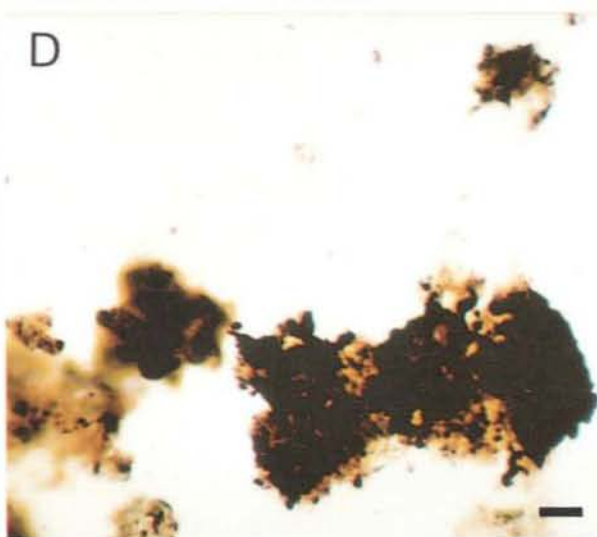
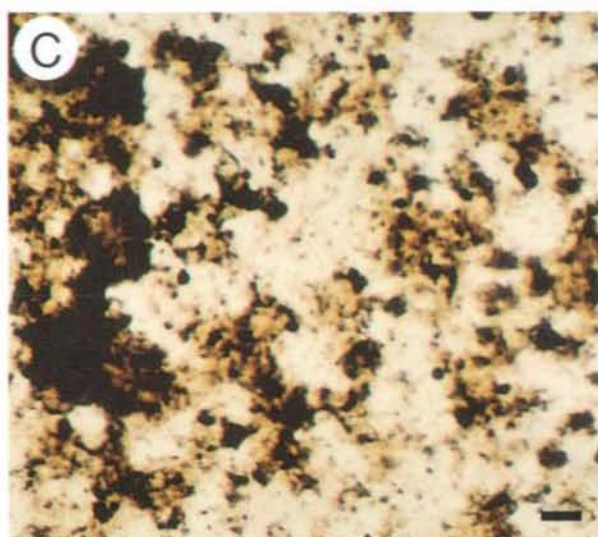
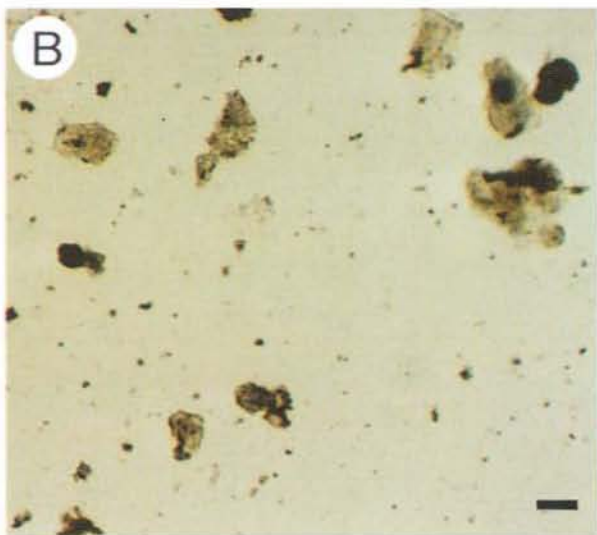
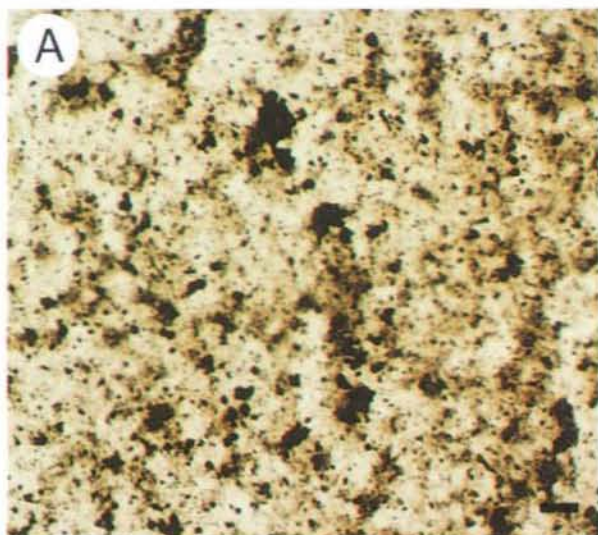




Plate 2. Cambrian palynomorphs

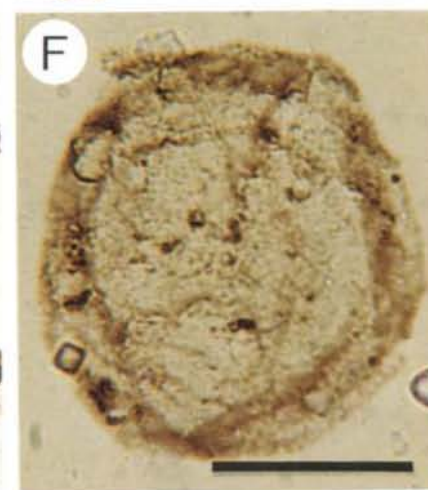
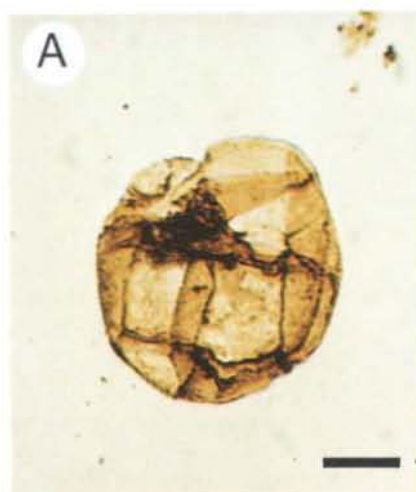


Plate 3. Ordovician palynomorphs

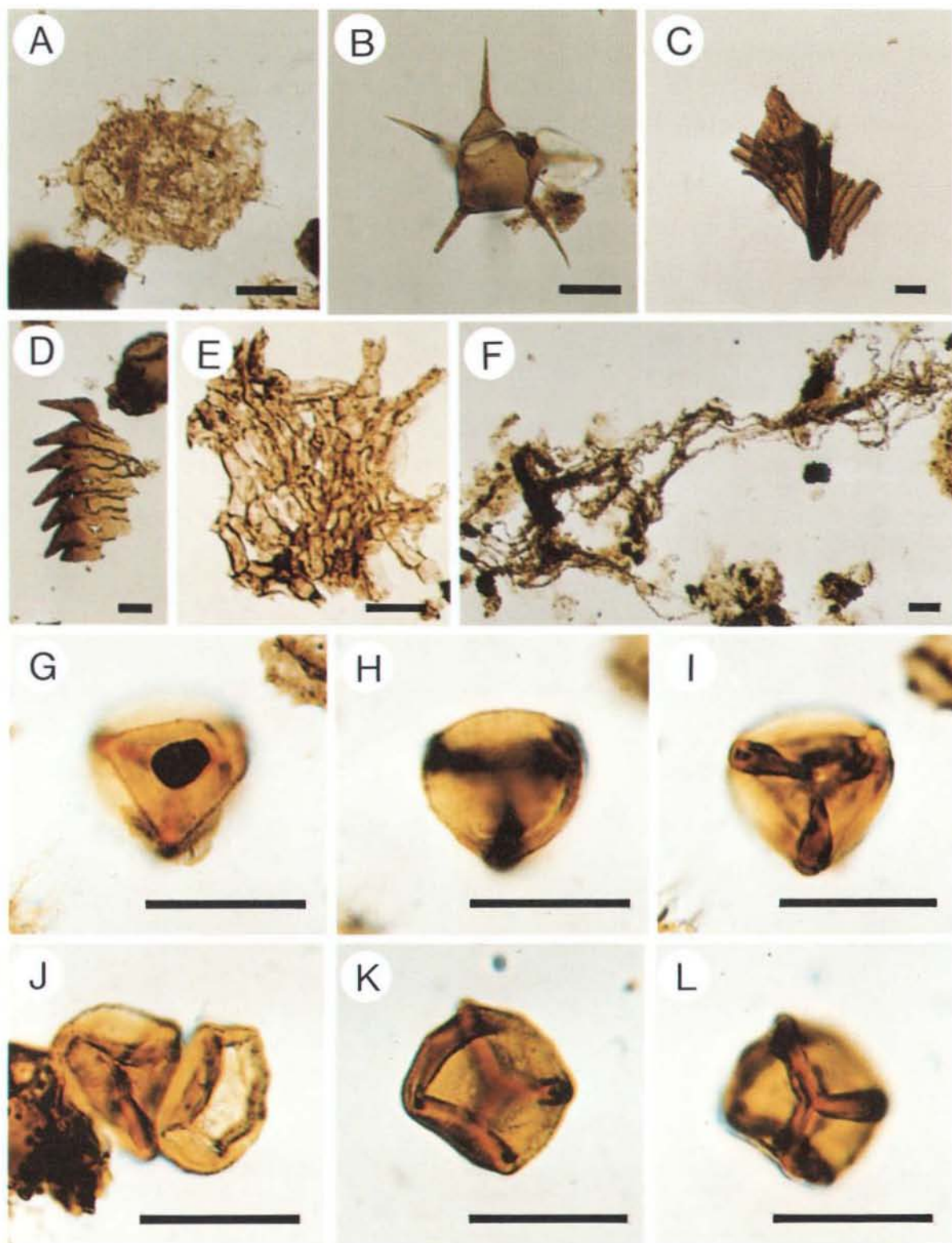
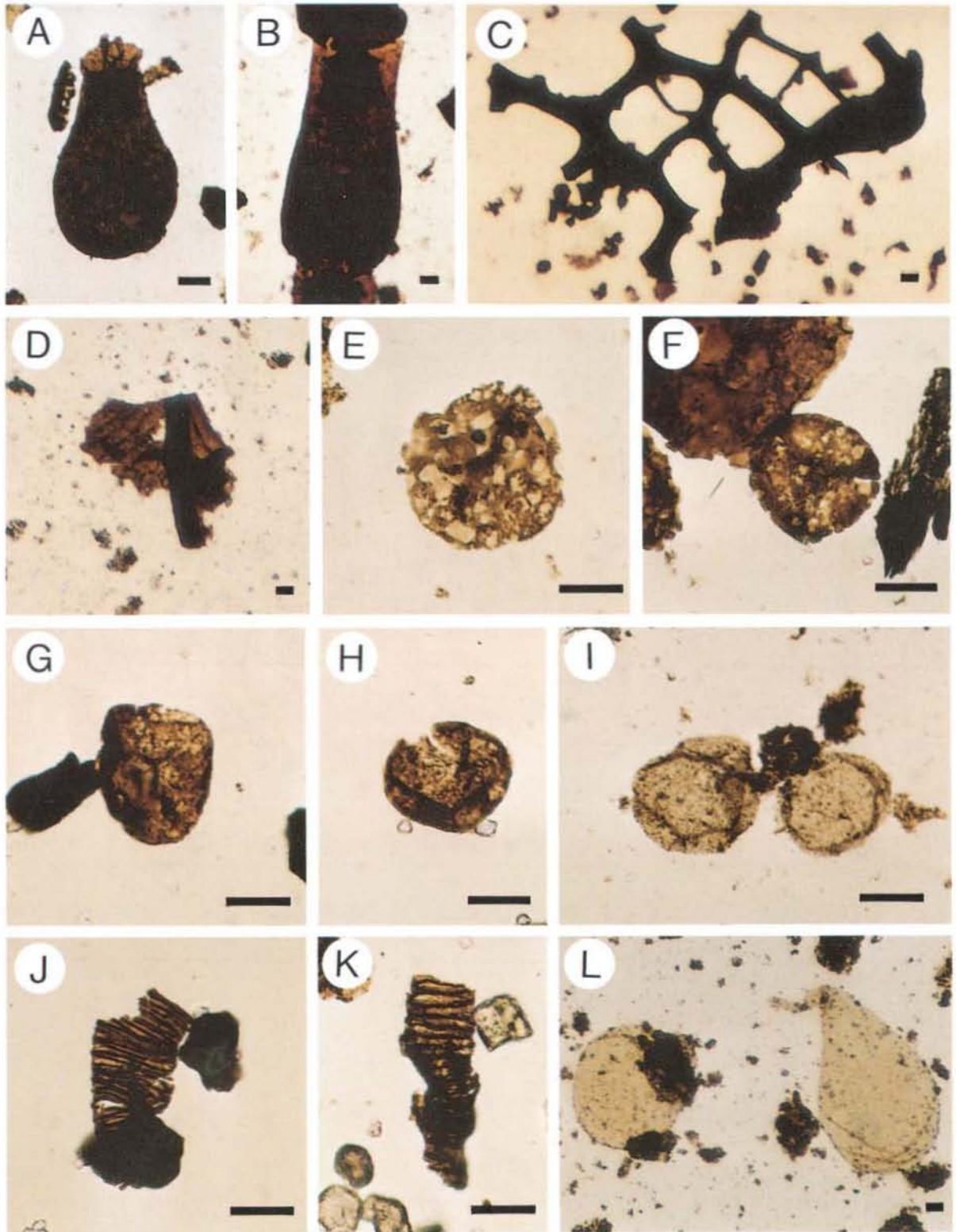


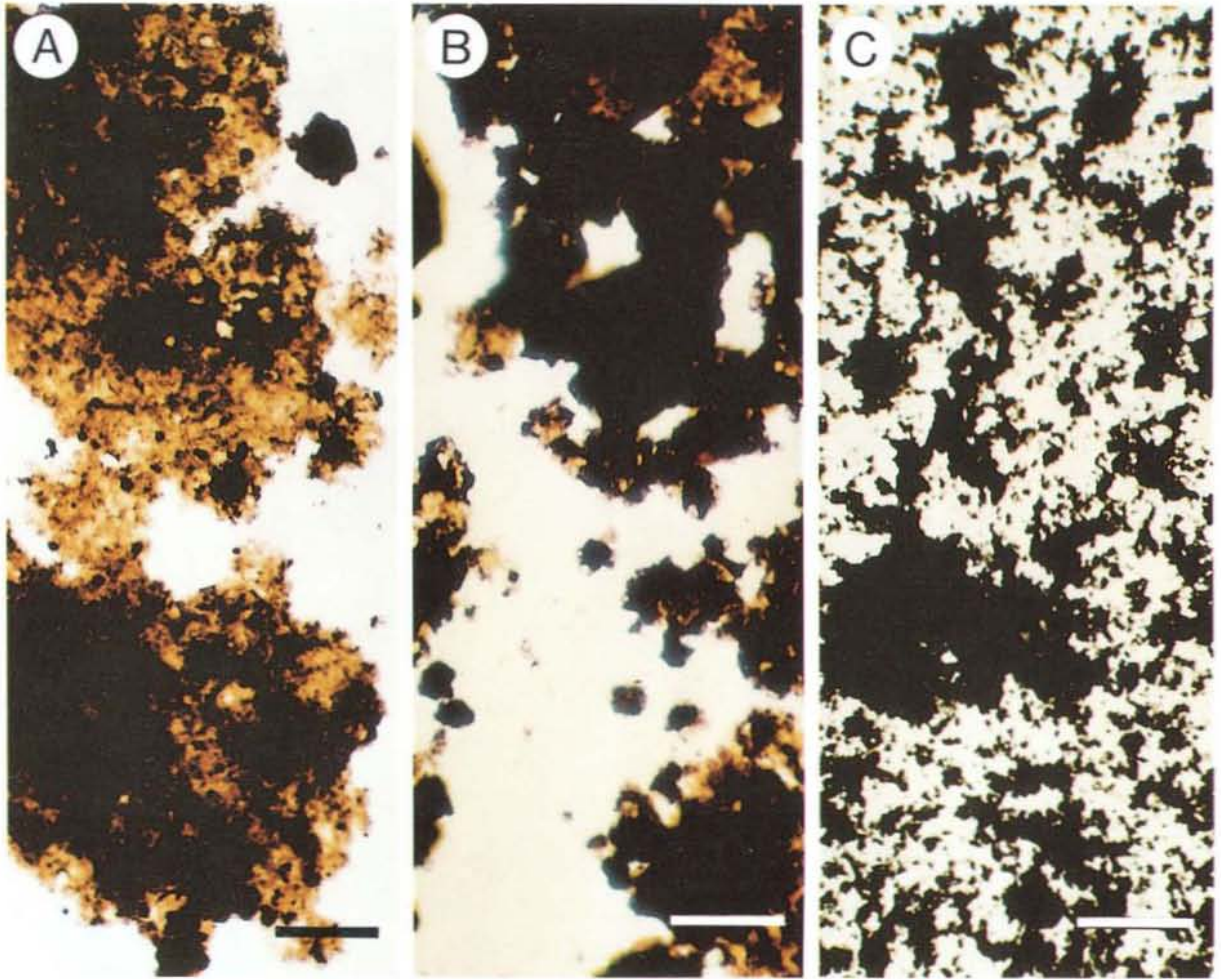


Plate 4. Silurian palynomorphs





**Plate 5.** Progressive coloration of amorphous kerogen with increasing thermal alteration

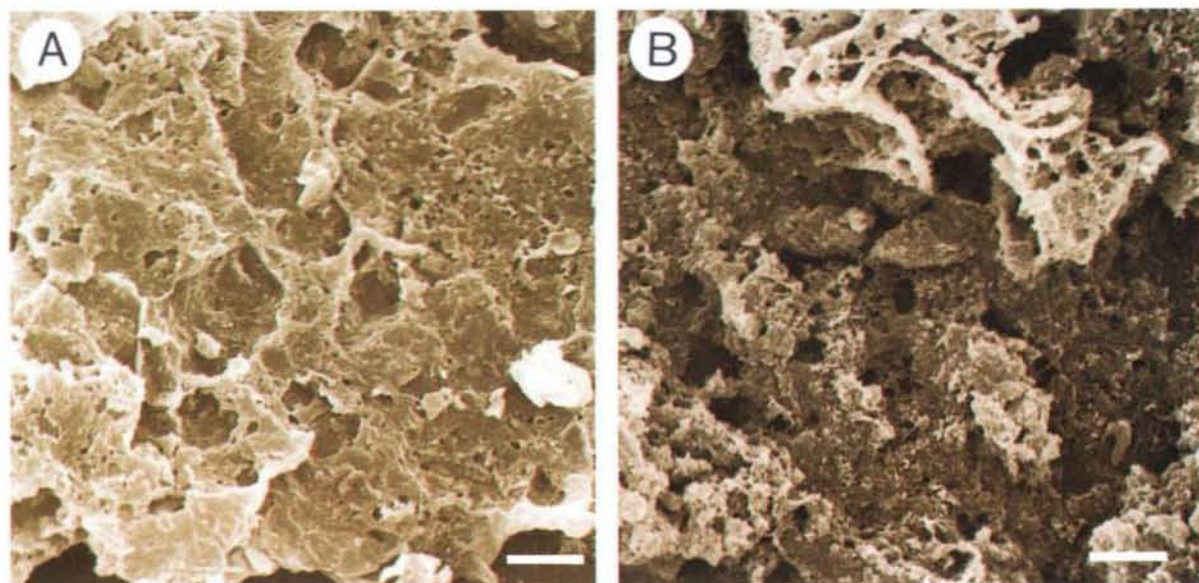


**Plate 5.** Progressive coloration of amorphous kerogen with increasing thermal alteration

Scale bar: 50  $\mu\text{m}$

A. TAI: (2)-2<sup>+</sup>, GGU 211759-2.    B. TAI: 2<sup>+</sup>-(3-), GGU 324405-2.    C. TAI: 4<sup>+</sup>, GGU 316475-1.

**Plate 6.** Change in structure of amorphous kerogen with increasing thermal alteration as observed in the scanning electron microscope



**Plate 6.** Change in structure of amorphous kerogen with increasing thermal alteration as observed in the scanning electron microscope

A. TAI: 2<sup>+</sup>-(3),  $T_{max}$ : 446, GGU 324405-2, scale bar: 10  $\mu$ m.

B. TAI: 4<sup>+</sup>,  $T_{max}$ : n.d., GGU 316475-2, scale bar: 10  $\mu$ m.

**Plate 7.** Field appearance of bitumen

A. Seeping asphalt from southern Wulff Land (equivalent to GGU 324200).

B. Hard solid bitumen in dolomite vug in the Sydpasset Formation (equivalent to GGU 324287-324299, core GGU 318003).

**Plate 8.** Macroscopic bitumen in slabs

A. Asphalt from seep in dolomite breccia. GGU 324200A.  
B. Like A. Stained. D0, D1, D2, C1, C2 correspond to generations of dolomite and calcite.

C. Asphalt from seep in dolomite breccia. Stained. Same generations of carbonates as A and B. GGU 324200E.

D. Hard solid bitumen in vugs and veins in dolomite grainstone. DO: dolomite grainstone, D1: saddle dolomite. Core GGU 318003.

**Plate 9.** Bitumen in thin section

A. Bitumen in coral. GGU 324130B, Lafayette Bugt Formation, Nyeboe Land. Plane light, stained, C: calcite, Fe-C: Fe-rich calcite, B: bitumen, scale bar: 1 mm.  
B. Bitumen in coral. GGU 316067, Lafayette Bugt Formation, Washington Land. Crossed nicols + gypsum plate. Q: quartz, C: calcite, B: bitumen, scale bar: 1 mm.  
C. Bitumen-filled fracture in calcarenite. GGU 318013-09, Lafayette Bugt Formation, Nyeboe Land. Plane light, scale bar: 2 mm.

D. Two-phased bitumen (black and yellow) (B1, B2) in saddle dolomite veins (D1) cross-cutting dolomite grainstone (D0). See close-up (arrow) in Plate 11. GGU 318003-53, Henson Gletscher Formation, Freuchen Land. Plane light, scale bar: 2 mm.

E.-F. Saddle dolomite vein (D1) in dolomite grainstone (D0). Bitumen occurs as impregnation in D0 (B1), as residual matter in the contact between D0 and D1 (B2,  $R_o$ : 0.92%) and in the centre of the vein (B3,  $R_o$ : 1.21%). GGU 318003-21, Sydpasset Formation, Freuchen Land. Plane light, scale bars: 1 mm.

**Plate 10.** Bitumen in palynologically prepared samples observed in microscope or in SEM

A. Bitumen with flaky appearance (note crystal impressions). GGU 315172-1, Ryder Gletscher Group Fm 6, Wulff Land. Scale bar: 25  $\mu$ m.  
B. Globular bitumen. GGU 315865-2, Aftenstjernesø Formation, Nares Land, scale bar: 50  $\mu$ m.  
C. Globular bitumen which has been extruded during and after sample preparation by the xylene-containing mounting medium. GGU 315199, Ryder Gletscher Group Fm 6, Warming Land. Scale bar: 12.5  $\mu$ m.

D. Bitumen with flaky appearance (note crystal impressions). GGU 315172-1, Ryder Gletscher Group Fm 6, Wulff Land. Scale bar: 10  $\mu$ m.

E. Bitumen mirroring imprints of crystals from coral space. GGU 316067-2, Lafayette Bugt Formation, Washington Land. Scale bar: 10  $\mu$ m.

F. Thread-like bitumen. GGU 324453, Buen Formation, Wulff Land. Scale bar: 25  $\mu$ m.



Plate 7. Field appearance of bitumen



Plate 8. Macroscopic bitumen in slabs

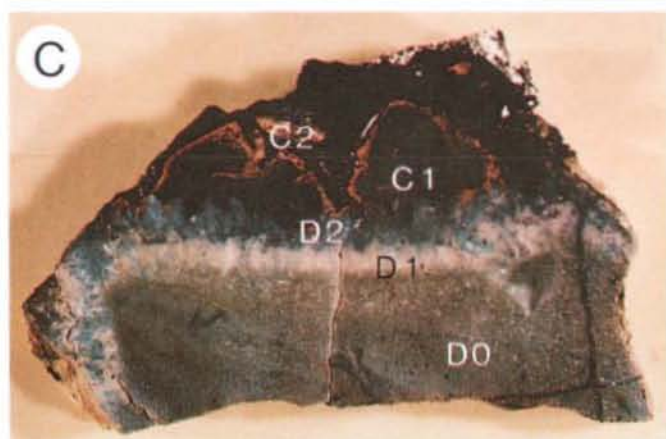
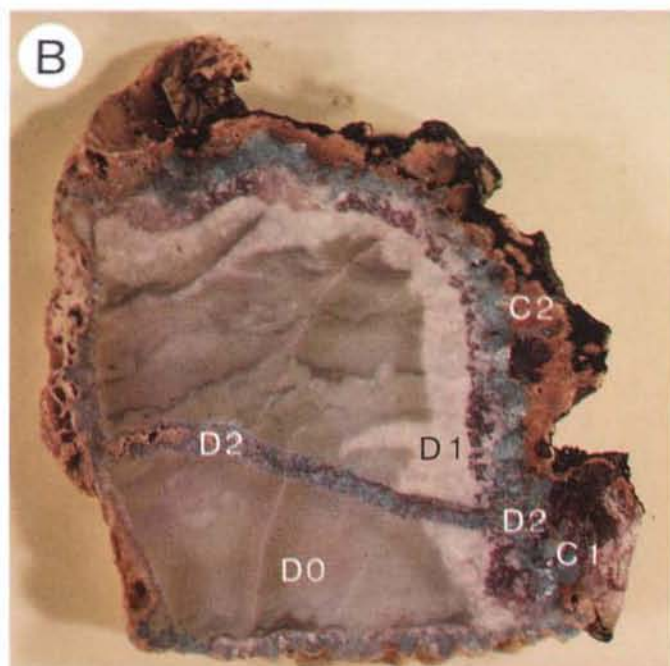
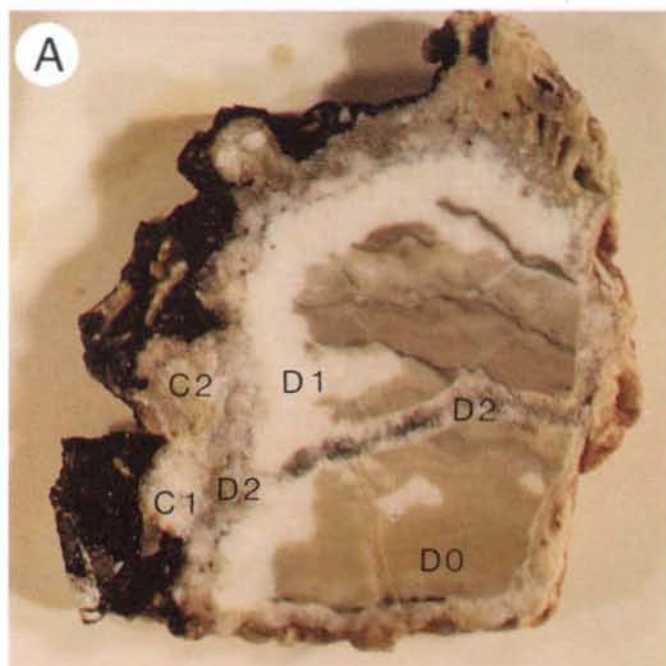




Plate 9. Bitumen in thin section

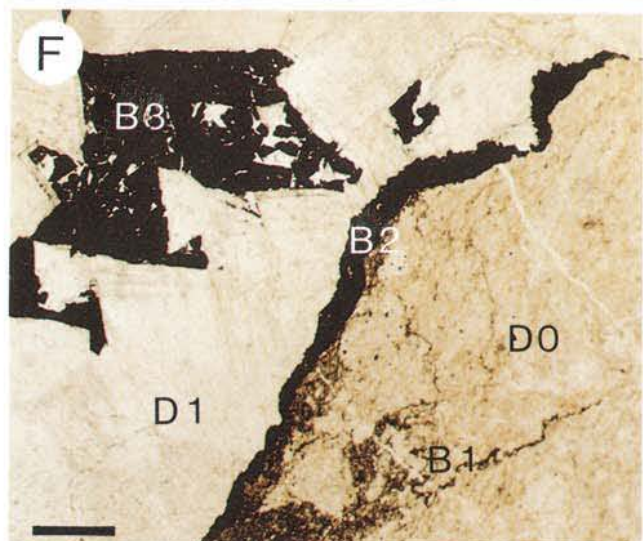
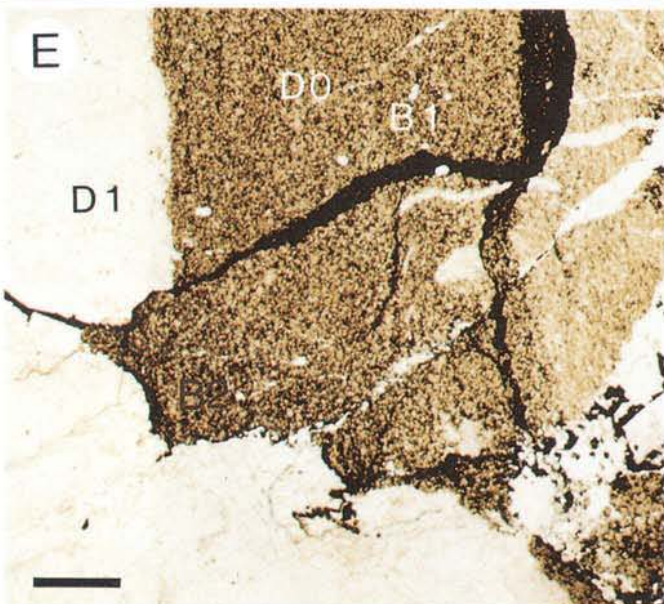
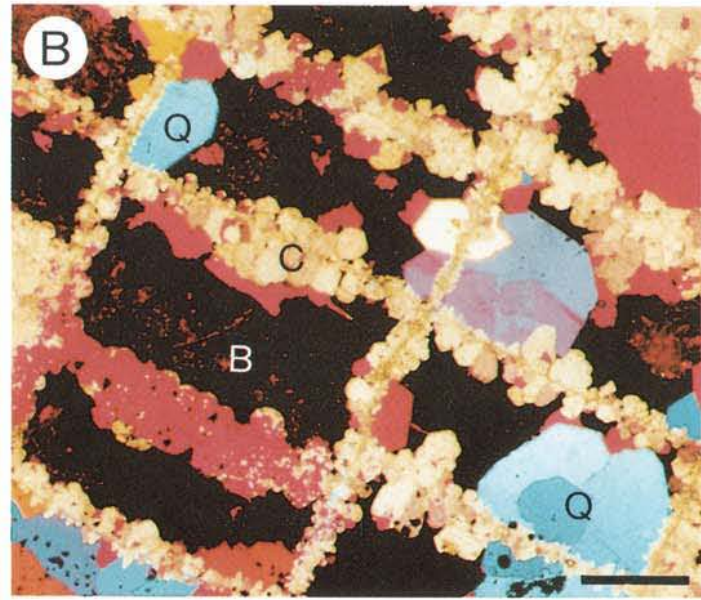
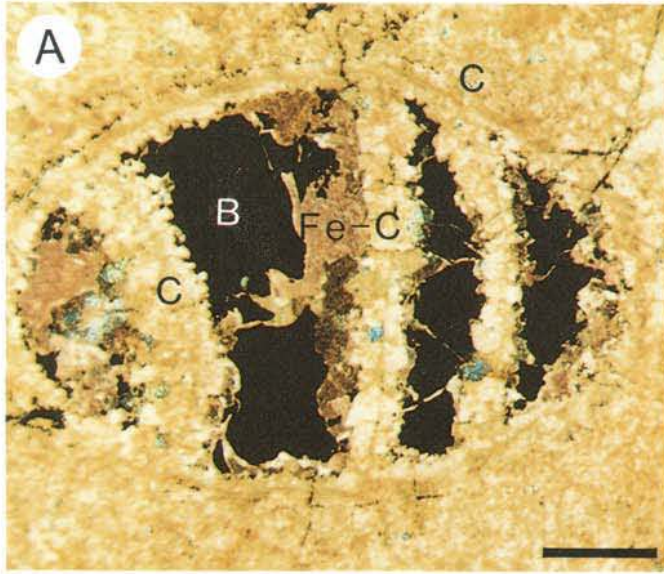




Plate 10. Bitumen in palynologically prepared samples observed in microscope or in SEM

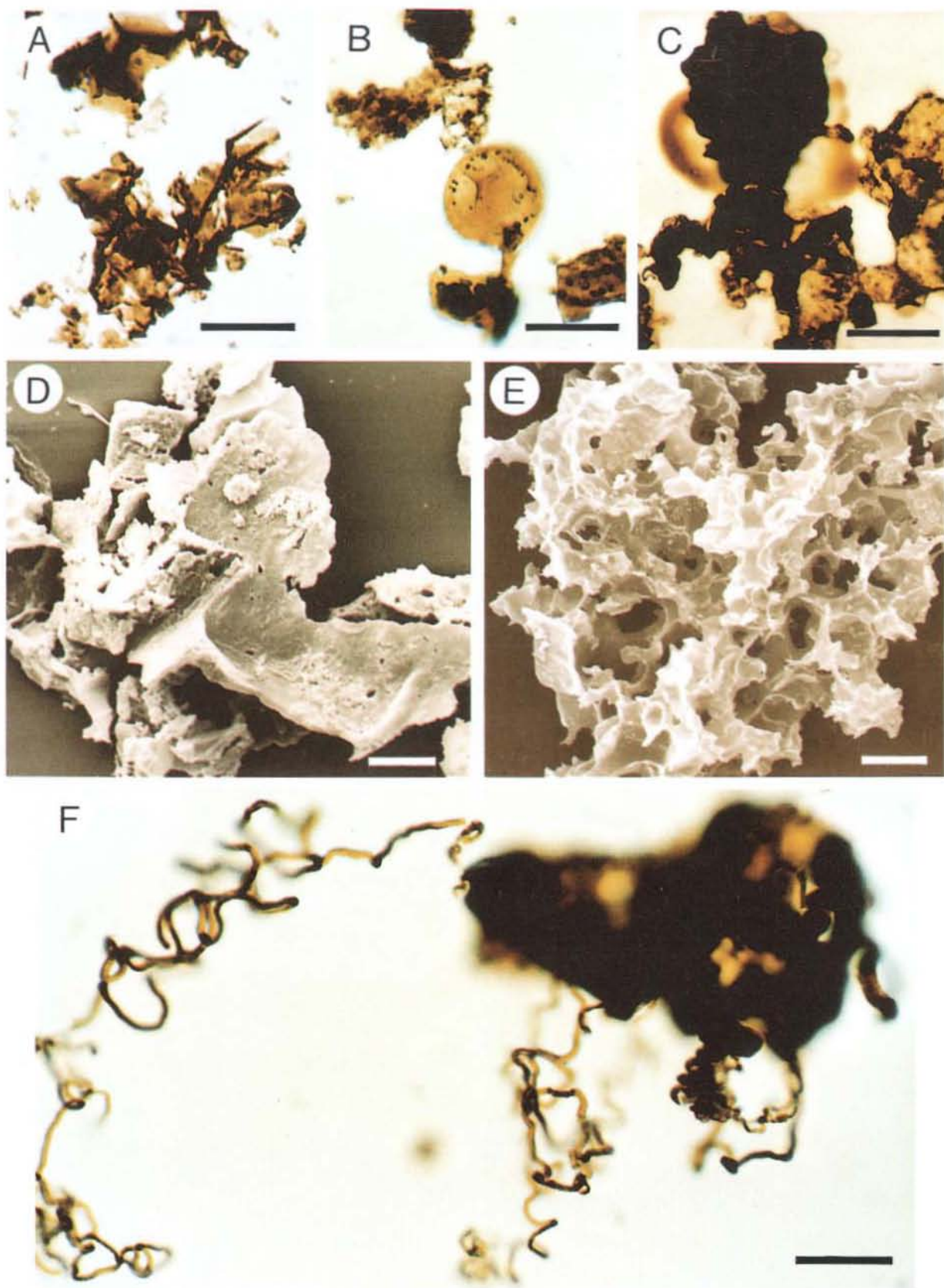


Plate 11. Bitumen in polished section

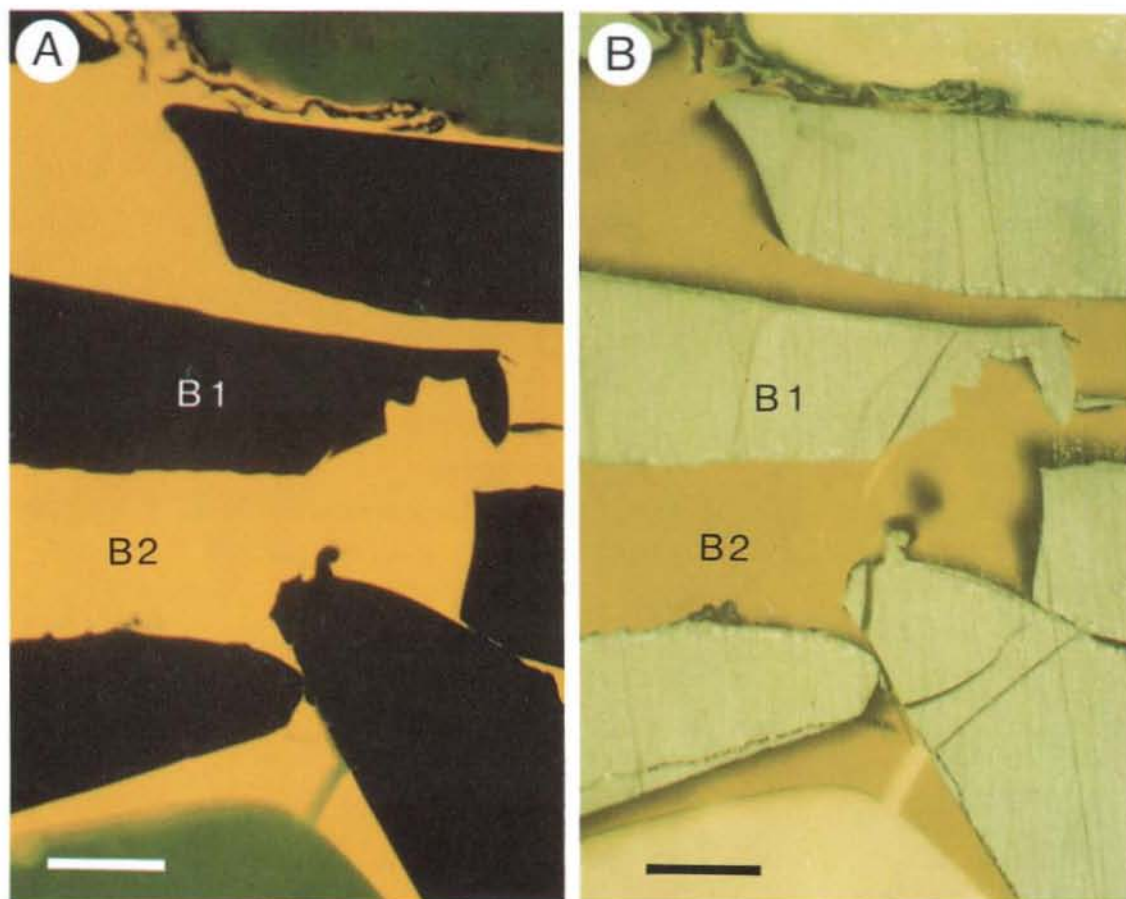


Plate 11. Bitumen in polished section

GGU 318003-53, Henson Gletscher Formation, Freuchen Land.

Scale bar: 50  $\mu\text{m}$

A. Fluorescent light photograph of two-phase bitumen.

B. Normal reflected light photograph of same field. The yellow-fluorescent low-reflecting bitumen (B2) has a  $R_o$  of 0.08% and the dark non-fluorescent high-reflecting bitumen (B1) a  $R_o$  of 1.17%.

*Cover picture*

Cambrian and Ordovician strata at Blue Cliffs, Wulff Land, North Greenland. Photo: J. Lautrup, GGU.